JFM RAPIDS
journals.cambridge.org/rapids

Signatures of fluid–fluid displacement in porous media: wettability, patterns and pressures

Bauyrzhan K. Primkulov[1], Amir A. Pahlavan[2], Xiaojing Fu[3], Benzhong Zhao[4], Christopher W. MacMinn[5] and Ruben Juanes[1],†

[1]Department of Civil and Environmental Engineering, Massachusetts Institute of Technology, Cambridge, MA 02139, USA

[2]Department of Mechanical and Aerospace Engineering, Princeton University, Princeton, NJ 08540, USA

[3]Department of Earth and Planetary Science, University of California at Berkeley, Berkeley, CA 94720, USA

[4]Department of Civil Engineering, McMaster University, Hamilton, ON, L8S 4L7, Canada

[5]Department of Engineering Science, University of Oxford, Oxford OX1 3PJ, UK

We develop a novel 'moving-capacitor' dynamic network model to simulate immiscible fluid–fluid displacement in porous media. Traditional network models approximate the pore geometry as a network of fixed resistors, directly analogous to an electrical circuit. Our model additionally captures the motion of individual fluid–fluid interfaces through the pore geometry by completing this analogy, representing interfaces as a set of moving capacitors. By incorporating pore-scale invasion events, the model reproduces, for the first time, both the displacement pattern and the injection-pressure signal under a wide range of capillary numbers and substrate wettabilities. We show that at high capillary numbers the invading patterns advance symmetrically through viscous fingers. In contrast, at low capillary numbers the flow is governed by the wettability-dependent fluid–fluid interactions with the pore structure. The signature of the transition between the two regimes manifests itself in the fluctuations of the injection-pressure signal.

Key words: porous media, fingering instability, capillary flows

1. Introduction

A beautiful array of flow patterns arises when a low-viscosity fluid displaces a more viscous fluid in a porous medium. The problem has been extensively examined through laboratory experiments, as well as numerical simulations and theoretical models (Saffman & Taylor 1958; Paterson 1981; Tryggvason & Aref 1983; Chen &

† Email address for correspondence: juanes@mit.edu

Wilkinson 1985; Kadanoff 1985; Måløy, Feder & Jøssang 1985; Nittmann, Daccord & Stanley 1985; Bensimon *et al.* 1986; Chen 1987; Homsy 1987; Arnéodo *et al.* 1989; Fernández *et al.* 1990; Li *et al.* 2009; Bischofberger, Ramachandran & Nagel 2015). The dynamics of such displacement can be characterized by two dimensionless groups: the ratio of viscous to capillary forces, or the capillary number (Ca), and the ratio of defending to invading fluid viscosities, or viscosity contrast (M). For high Ca, the resulting displacement patterns are reminiscent of diffusion-limited aggregation (Witten, Sander & Sander 1981; Niemeyer, Pietronero & Wiesmann 1984; Daccord, Nittmann & Stanley 1986; Meakin, Tolman & Blumen 1989; Conti & Marconi 2010). For low Ca, the displacement dynamics becomes more intricate, and the emerging patterns display a strong dependence on the pore geometry (Lenormand, Zarcone & Sarr 1983; Lenormand & Zarcone 1985; Lenormand, Touboul & Zarcone 1988; Fernandez, Rangel & Rivero 1991; Måløy *et al.* 1992; Furuberg, Måløy & Feder 1996; Ferer *et al.* 2004; Toussaint *et al.* 2005; Holtzman, Szulczewski & Juanes 2012) and the wettability of the medium – that is, the chemical affinity of the solid for each fluid (Stokes *et al.* 1986; Trojer, Szulczewski & Juanes 2015; Zhao, MacMinn & Juanes 2016; Odier *et al.* 2017). In particular, an intermittent injection-pressure signal emerges in the limit of low Ca (Måløy *et al.* 1992; Furuberg *et al.* 1996). Given that in most practical applications visualization of the flow in porous media is not possible, the pressure signal is often the only source of information. Surprisingly, no modelling approach to date has been able to capture the injection-pressure signal across different Ca and pore wettabilities. Here, we develop a new pore-network model that fills this gap, and we use it to explore the transition from viscous-dominated to capillary-dominated flow regimes by examining the connections among fluid morphology and pressure signal.

Pore-network models of flow in porous media can be broadly classified into two groups: quasi-static and dynamic models (Blunt 2001; Meakin & Tartakovsky 2009; Joekar-Niasar & Hassanizadeh 2012). Quasi-static models neglect viscous effects and advance the invading fluid through either invasion-percolation (Chandler *et al.* 1982; Lenormand *et al.* 1988) or event-based algorithms (Cieplak & Robbins 1988, 1990). Although a quasi-static approach can be effective in reproducing experimental invasion patterns at low Ca (Primkulov *et al.* 2018), it is unable to capture the temporal evolution of the injection-pressure signal. Dynamic network models approximate the flow channels with a network of interconnected capillary tubes. Viscous pressure drops are calculated by assuming fully developed viscous flow within each tube. Local capillary pressures within the network are calculated from either the interface position within pore throats (Aker *et al.* 1998*b*; Gjennestad *et al.* 2018) or through mass balance of the two phases in pore bodies (Al-Gharbi & Blunt 2005; Joekar-Niasar, Hassanizadeh & Dahle 2010). Another notable class of models is invasion percolation in a gradient: a percolation model designed to incorporate buoyancy forces (Wilkinson 1984; Birovljev *et al.* 1991; Frette *et al.* 1992; Meakin *et al.* 1992), and then extended to model (linear) pressure gradients (Yortsos, Xu & Salin 1997). None of the studies of invasion percolation in a gradient, however, incorporate any notion of wettability (they all deal exclusively with strong drainage), pore-scale dynamics, or capillary-number-dependent pressure fluctuations.

In fact, most existing pore-network models, both quasi-static and dynamic, are limited to strong drainage (or injection of non-wetting fluid) and do not include wettability-induced cooperative pore filling (Aker *et al.* 1998*b*; Al-Gharbi & Blunt 2005; Holtzman & Juanes 2010; Joekar-Niasar *et al.* 2010). The only dynamic pore-network model to date that includes cooperative pore-filling events (Holtzman &
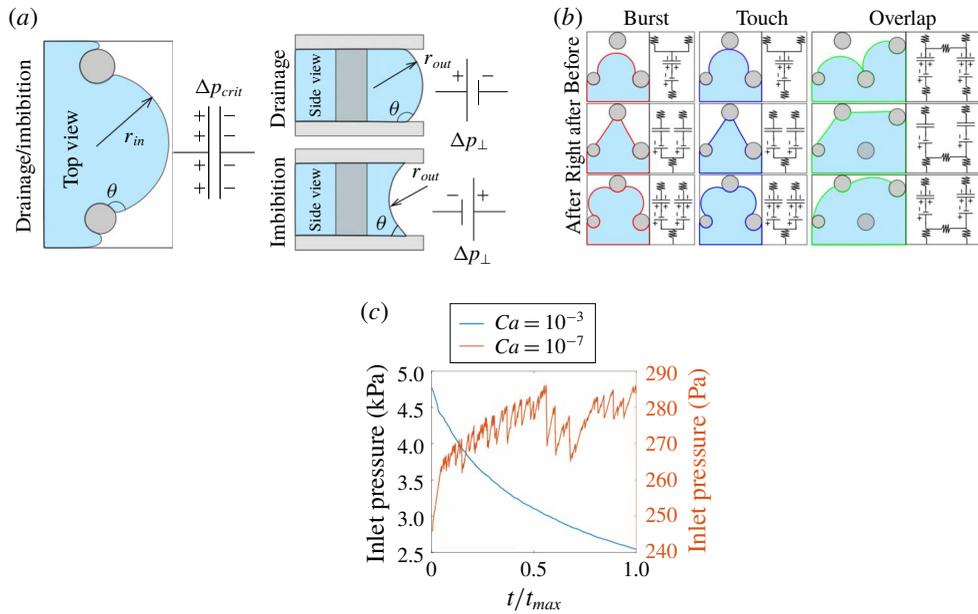
FIGURE 1. (*a*) Schematic diagram of in-plane and out-of-plane curvatures within the flow cell. Out-of-plane curvature represents the overall affinity of the porous medium to the invading fluid. It is determined by θ and is analogous to a battery. In-plane curvature changes as the local interface evolves while pinned to a pore throat, and it is analogous to a capacitor. (*b*) Evolution of burst, touch, and overlap events. (*c*) Temporal profiles of the injection pressure bear close resemblance to similar experiments in the drainage regime at low (orange) and high (blue) *Ca* (Furuberg *et al.* 1996; Zhao *et al.* 2016).

Segre 2015) does so by combining pore-level invasion events of Cieplak & Robbins (1988, 1990) with viscous relaxation through the pore network. This viscous-relaxation assumption is at odds with the physics of interface motion in the capillary-dominated regime and, as a result, this model is unable to capture the injection-pressure signal observed experimentally in the limit of intermediate and low *Ca* (Måløy *et al.* 1992; Furuberg *et al.* 1996; Zhao *et al.* 2016). We present in § 2 a consistent framework that combines viscous, capillary, and wettability effects in a single dynamic network model that builds a direct analogy between local fluid–fluid interfaces and electric capacitors. Our model reproduces, quantitatively, the fluid–fluid displacement patterns for a wide range of *Ca* and wettabilities (§ 3), and points to a surprising and heretofore unrecognized transition in the pressure fluctuations between the low- and high-*Ca* flow regimes (§ 4).

2. Moving-capacitor model

Consider a moving fluid–fluid interface in a micromodel (figure 1*a*). Neglecting dynamic-contact-angle effects (Hoffman 1975) for simplicity, the shape of the meniscus between posts is uniquely defined by the combination of Laplace pressure and substrate wettability, defined through a contact angle θ at which the interface meets post surfaces (Cieplak & Robbins 1988, 1990). As the interface advances, the Laplace pressure increases until the interface encounters a burst, touch or overlap event, as defined by Cieplak & Robbins (1988, 1990). The burst event is equivalent

to a Haines jump (Haines 1930; Berg *et al.* 2013), while the touch and overlap events take place when the local interface either touches the nearest opposing post or coalesces with a neighbouring interface, respectively (figure 1*b*). If the interface becomes unstable due to burst or touch, a single pore is invaded and two new interfaces appear. In the case of an overlap event, two (in some cases more) pores are filled simultaneously. These pore-level events are an integral part of the model and, indeed, this sensitivity is what permits capturing wettability effects within the model. The events evolve differently at different wettabilities – burst events are most frequent in drainage, while touch and overlap are most frequent in imbibition (or injection of wetting fluid) (Cieplak & Robbins 1990; Primkulov *et al.* 2018).

We can explicitly calculate the critical Laplace pressure Δp_{crit} corresponding to all events from the values of the contact angle, radii and coordinates of the posts (Primkulov *et al.* 2018), and thus can use the analogy between electric capacitors and fluid–fluid interfaces in constructing our network model. A capacitor represents the pinning of the fluid–fluid interface at a pore throat, and is active in both drainage and imbibition: the interface moves only when a local depinning threshold (Δp_{crit}) is reached, and the fluid front moves to restart the pinning–depinning cycle from zero in-plane curvature (figure 1*b*). This progression of the in-plane curvature in our model was motivated by the work of Cieplak and Robbins (Cieplak & Robbins 1988, 1990) (see also Rabbani *et al.* 2018) and experiments on the progression of the in-plane curvature between the Hele-Shaw cell posts (Jung *et al.* 2016; Lee *et al.* 2017). This is what allows capturing pressure fluctuations in the limit of low *Ca* (figure 1*c*). The battery analogy represents the overall affinity of the porous medium to the invading fluid, set by the out-of-plane curvature at the fluid front. The out-of-plane curvature is fixed throughout a single simulation, and determined by the value of the contact angle (given the constant gap between the flow-cell plates): it is positive in drainage and negative in imbibition (figure 1*a*). To complete the analogy between an electric circuit and a pore network, one can think of a network of resistors being responsible for viscous effects, capacitors and batteries responsible for capillary effects, and local rules for circuit rearrangements responsible for wettability effects (figure 1*b*).

Therefore, the pressure drop across an edge of the network containing a fluid–fluid interface has three components: (i) pressure drop due to viscous dissipation, (ii) Laplace pressure drop due to in-plane curvature of the interface, and (iii) Laplace pressure drop due to out-of-plane curvature of the interface. We calculate the viscous pressure drop assuming Poiseuille flow in a capillary tube, which is analogous to the potential drop across a resistor. The out-of-plane component of the Laplace pressure can be expressed as either a positive or negative pressure jump ($\Delta p_{\perp} = -(2\gamma \cos \theta / h)$, where γ is the interfacial tension, and h is the cell height) depending on the substrate wettability; this is analogous to a battery in an electric circuit. The Laplace pressure due to in-plane curvature of the interface is analogous to a capacitor which allows flow until it reaches the critical pressure ($\Delta p_{crit} = \min\{p_{burst}, p_{touch}, p_{overlap}\}$). Since we can calculate Δp_{crit} for all edges at the invading fluid front, we use a linear estimate of the in-plane Laplace pressure drops within our network ($\Phi(t)\Delta p_{crit}$), where $\Phi(t)$ stands for the filling ratio of a given throat. When $\Phi(t) \to 0$, the in-plane Laplace pressure is negligible. When $\Phi(t) \to 1$, the throat is nearly full and has a critical in-plane Laplace pressure Δp_{crit}. This analogy between local interfaces and capacitors allows us to incorporate local changes in Laplace pressure due to filling of pore throats. Once a node in the network reaches its maximal potential, which coincides with its filling capacity, it becomes unstable and the interface advances. We assume that the in-plane and out-of-plane Laplace pressures are decoupled, and this

is done to maintain the simplicity of the overall model. With this assumption, one can run the model for either $h/a \gg 1$ or $h/a \ll 1$, where these conditions would result in negligible or dominant contributions of the out-of-plane curvature in the model, respectively.

The topology of the pore network is captured through the incidence matrix $\boldsymbol{A}$ by examining the adjacency of the pores (Strang 2007). We number all pores and adopt the convention that pore connections are oriented in the direction of increasing pore numbers. Rows of $\boldsymbol{A}$ represent edges, and columns of $\boldsymbol{A}$ represent nodes of the network. We also make use of the diagonal conductance matrix $\boldsymbol{C}$, whose elements are hydraulic conductivities of the network edges. The elements of this matrix can be calculated as $c = \pi r^4 / 8\mu L$, assuming fully developed Hagen–Poiseuille flow through a rectangular tube with hydraulic radius r and length L, where μ is the effective viscosity of the fluid in the channel.

The pressure difference across the network edges can be calculated as $\boldsymbol{e} = \boldsymbol{b} - \boldsymbol{Ap}$, where $\boldsymbol{b}$ and $\boldsymbol{p}$ stand for pressure change due to out-of-plane contribution to Laplace pressure (batteries) and node pressures, respectively. The network flow rates can be calculated from this pressure difference as $\boldsymbol{q} = \boldsymbol{Ce}$. At the same time, flow rates must obey mass conservation, $\boldsymbol{A}^T \boldsymbol{q} = \boldsymbol{f}$, where $\boldsymbol{f}$ stands for flow sources at the nodes. After eliminating $\boldsymbol{e}$, the flow through the network without the in-plane contribution to Laplace pressure (capacitors) is obtained through the following system of equations:

$$\boldsymbol{q} = \boldsymbol{C}(\boldsymbol{b} - \boldsymbol{Ap}), \tag{2.1}$$

$$\boldsymbol{A}^T \boldsymbol{q} = \boldsymbol{f}. \tag{2.2}$$

We set constant flow boundary conditions at the inlet pores (at the centre of the flow cell) and constant-pressure boundary conditions at the outlet pores (at the edges of the flow cell). We note that $\boldsymbol{Ap}$ can be decomposed into components of nodes with prescribed pressure and all other nodes ($\boldsymbol{Ap} = \boldsymbol{A}_{outer}\boldsymbol{p}_{outer} + \tilde{\boldsymbol{A}}\tilde{\boldsymbol{p}}$), and therefore (2.1)–(2.2) transform to

$$\begin{bmatrix} \boldsymbol{C}^{-1} & \tilde{\boldsymbol{A}} \\ \tilde{\boldsymbol{A}}^T & \boldsymbol{0} \end{bmatrix} \begin{bmatrix} \boldsymbol{q} \\ \tilde{\boldsymbol{p}} \end{bmatrix} = \begin{bmatrix} \boldsymbol{b} - \boldsymbol{A}_{outer}\boldsymbol{p}_{outer} \\ \tilde{\boldsymbol{f}} \end{bmatrix} = \begin{bmatrix} \tilde{\boldsymbol{b}} \\ \tilde{\boldsymbol{f}} \end{bmatrix}. \tag{2.3}$$

The solution to (2.3) provides values of both edge flow rates and node pressures for given boundary conditions.

Finally, we incorporate the pressure drop due to in-plane Laplace pressure (capacitors) within the network. Taking into account the direction of the edges (an array $\boldsymbol{d}(t)$ consisting of 1 and -1), the total pressure drop across the network edges can be written as $\boldsymbol{e} = \tilde{\boldsymbol{b}} - \tilde{\boldsymbol{A}}\tilde{\boldsymbol{p}} - \boldsymbol{d}(t)\Phi(t)\Delta p_{crit}$. In other words, the in-plane Laplace pressure is the product of the filling ratio and the critical pressure from the quasi-static model (Primkulov *et al.* 2018). Therefore, the equations governing two-phase flow through the network can be written as

$$\begin{bmatrix} \boldsymbol{C}^{-1}(t) & \tilde{\boldsymbol{A}} \\ \tilde{\boldsymbol{A}}^T & \boldsymbol{0} \end{bmatrix} \begin{bmatrix} \boldsymbol{q}(t) \\ \tilde{\boldsymbol{p}}(t) \end{bmatrix} = \begin{bmatrix} \tilde{\boldsymbol{b}} - \boldsymbol{d}(t)\Phi(t)\Delta p_{crit} \\ \tilde{\boldsymbol{f}} \end{bmatrix}. \tag{2.4}$$

We now discuss the mechanics of the time stepping in our two-phase flow model. After we initialize the interface locations within the network, we use an adaptive forward Euler time stepping to update the filling ratios of the network edges at the

interface $\Phi(t)$. We ensure that only a fraction of the edge total volume at the interface flows within the time step (Aker *et al.* 1998*b*). After every time step, we use $\Phi(t)$ to update the conductance matrix $\mathbf{C}(t)$ and resolve the flow through (2.4) with updated pressure drops across the fluid–fluid front.

In the spirit of the fundamental contributions from Cieplak & Robbins (1988, 1990), our model takes the form of an arrangement of cylindrical posts confined between the plates of a Hele-Shaw cell. The approach is simple enough to lead to universal findings, yet sufficiently complex to have direct relevance to microfluidic geometries, as well as engineered and natural porous media – much like Lenormand's phase diagram (Lenormand *et al.* 1988). By doing so, we demonstrate the ability to reproduce physics – in particular, pressure fluctuations under a wide range of wetting conditions – which, until now, were inaccessible to pore-network modelling. A limitation of the model presented here is that it does not extend to contact angles below 45°, where the wetting fluid preferentially wets the corners of the pore geometry at low *Ca* and forms film flow at high *Ca* (Zhao *et al.* 2016; Odier *et al.* 2017).

3. Invasion patterns

We simulate immiscible fluid–fluid displacement by setting a constant injection rate at the centre of the flow cell and zero pressure at the outlets. The invading and defending fluid viscosities are set to 8.9×10^{-4} Pa s and 0.34 Pa s respectively. The post height h is 100 µm, and interfacial tension γ is set to 13×10^{-3} N m^{-1}. These parameters as well as the pore geometry are chosen to mimic the experiments of Zhao *et al.* (2016). The flow cell has an outer diameter of 30 cm. We perform simulations for wetting conditions from strong drainage ($\theta = 160°$) to weak imbibition ($\theta = 46°$). Figure 1(*c*) shows the pressure profiles for $\theta = 160°$ at $Ca \in \{10^{-3}, 10^{-7}\}$, respectively. In the limit of high *Ca*, the more viscous defending fluid sustains substantial spatial pressure gradients, and the injection pressure gradually drops as more of the defending fluid is displaced (Zhao *et al.* 2016). In contrast, in the limit of low *Ca*, the pressure field is virtually uniform in each fluid, and the injection pressure exhibits intermittent fluctuations typical of slow capillary-dominated drainage (Måløy *et al.* 1992; Aker, Måløy & Hansen 1998*a*; Knudsen & Hansen 2002; Moebius & Or 2012).

The morphology of the invading fluid at breakthrough can be analysed by means of a binary-image representation of the invasion patterns (Cieplak & Robbins 1988, 1990; Primkulov *et al.* 2018) (figure 2*a*). We estimate the width and number of fingers in the invading fluid pattern following the protocol outlined in Cieplak & Robbins (1988, 1990) and modified in Primkulov *et al.* (2018). The binary image is sliced horizontally and vertically, with each slice containing clusters of invading fluid pixels. We calculate the finger width as the mean size of these clusters. Figure 2(*d*) shows that the finger width, normalized by the typical pore size, increases as $\theta \to 46°$ for all *Ca*, which is in agreement with experimental observations (Stokes *et al.* 1986; Trojer *et al.* 2015; Zhao *et al.* 2016). While figure 2(*a*) demonstrates that the number of fingers increases with *Ca* (Lenormand *et al.* 1988; Fernández *et al.* 1990; Zhao *et al.* 2016), we observe an unexpected behaviour (figure 2*b*): the finger density changes with the substrate wettability, and exhibits a maximum at approximately $\theta = 90°$. This effect is most pronounced for $10^{-6} < Ca < 10^{-3}$ (when viscous and capillary effects are comparable).

We explain the peak in the viscous finger density at $\theta \approx 90°$ in figure 2(*b*) by considering in-plane and out-of-plane contributions to the Laplace pressure. At a fixed *Ca*, the ratio of viscous and capillary forces in the micromodel changes as a function
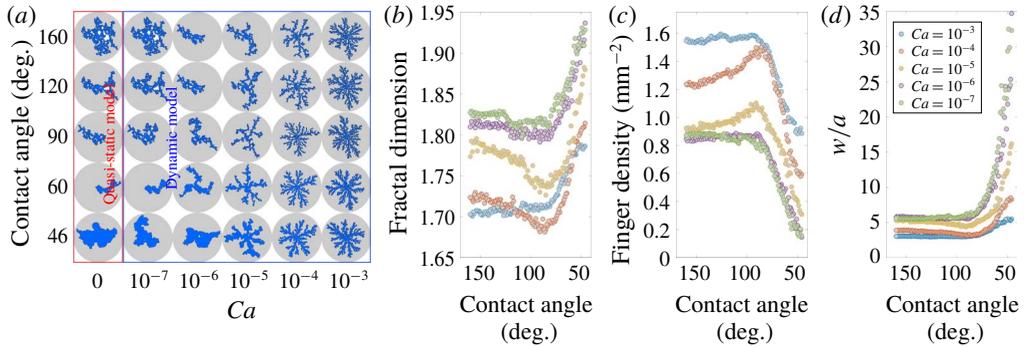
FIGURE 2. (a) Phase diagram of the invading fluid morphology at breakthrough; (b) fractal dimension, computed by means of the box-counting method; (c) number of fingers per unit area of injected fluid, which exhibits a maximum near $\theta = 90°$; (d) normalized finger width (w/a) at different Ca and wettabilities measured at breakthrough. Finger width increases as the posts become more wetting to the invading fluid.

of substrate wettability. The capillary forces have out-of-plane contributions, which are nominally equal to zero when $\theta = 90°$, so the ratio of viscous and capillary forces increases as θ changes from 160° to 90° at fixed Ca. In addition, when θ changes from 90° to 46°, the cooperative pore-filling mechanisms become dominant and widen the largest fingers, which in turn consume the smaller ones and reduce the number of fingers. The combination of these two effects results in the local maximum in the number of viscous fingers around $\theta \approx 90°$ across different Ca (figure 2b).

For a contact angle θ near 160° (strong drainage) and high values of Ca (10^{-3} and 10^{-4}), the invading fluid front advances through viscous fingers with fractal dimension close to 1.71, typical of diffusion-limited aggregation (DLA)-type morphology (Witten *et al.* 1981). As Ca is reduced to a low value (10^{-7}), the fractal dimension increases to approximately 1.82, characteristic of invasion percolation (Wilkinson & Willemsen 1983) (figure 2b). This increasing trend in fractal dimension is consistent with the decrease in finger density (figure 2c) and the increase in finger width (figure 2d).

As the contact angle approaches 46°, cooperative pore filling becomes the dominant flow mechanism at all values of Ca. This flow regime results in the compact displacement of the defending fluid, and thus the fractal dimension increases, approaching a value of 2 at low Ca, indicative of stable displacement.

4. Pressure signature

The fundamental difference in the fluid–fluid displacement process between low and high Ca is reflected in the temporal injection-pressure signals (figure 3). When the capillary number is relatively high ($Ca = 10^{-3}$), viscous forces dominate, and the injection pressure decreases with time for all substrate wettabilities (Zhao *et al.* 2016) (figure 3a). Here, most of the pressure drop takes place in the more viscous defending fluid. Consequently, as more of the defending fluid is displaced, the pressure required to maintain the prescribed injection flow rate decreases. In contrast, at $Ca = 10^{-7}$, viscous dissipation is negligible, and the injection pressure is determined by the sum of outlet and Laplace pressures. As a result, the injection pressure fluctuates in a stick–slip manner around a mean value (figure 3b), as
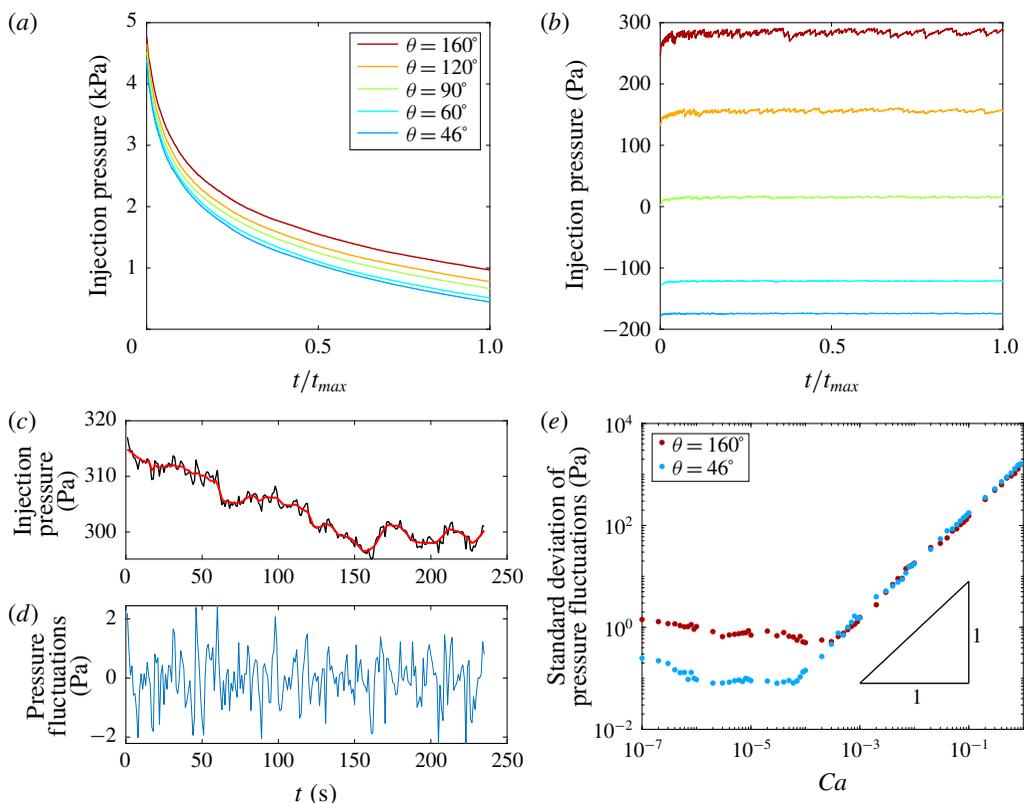
FIGURE 3. (*a,b*) Temporal evolution of the injection pressure at $Ca = 10^{-3}$ and $Ca = 10^{-7}$ respectively. At high Ca, the injection pressure decreases as the viscous fingers approach the outer boundary of the flow cell. At low Ca, the injection pressure is dominated by Laplace pressure fluctuations at the interface. We use wavelet decomposition (Cai 2002; Sygouni, Tsakiroglou & Payatakes 2006, 2007) to split the pressure signal ($Ca = 10^{-5}$ and $\theta = 160°$ here) into its (*c*) global trend and (*d*) cyclic component. (*e*) The standard deviation of the pressure fluctuations point at two different regimes. At low Ca, pressure fluctuations are dominated by stick–slip changes in Laplace pressure. At high Ca, pressure fluctuations are dominated by changes in the effective hydraulic conductance of dominant flow channels.

has been documented in slow drainage experiments (Måløy *et al.* 1992; Furuberg *et al.* 1996; Moebius & Or 2012). The pressure signals in figure 3(*b*) highlight the roles that in-plane and out-of-plane curvatures play in our model. Out-of-plane curvature plays the role of batteries, and thus provides additional resistance/drive (in drainage/imbibition, respectively) to the flow at the interface. The magnitude of the pressure drop/rise at the batteries is a function of wettability, which explains why the mean value of the injection-pressure signal also varies with wettability (figure 3*b*). The in-plane curvature plays the role of capacitors. As the invading fluid is injected, the in-plane component of Laplace pressure grows at the interface until the meniscus near the pore with lowest critical entry pressure becomes unstable due to burst, touch or overlap. This results in the rapid advance of the local interface, which pressurizes the defending fluid ahead. This overpressure then dissipates (see movie S1 in the supplementary materials available at https://doi.org/10.1017/jfm.2019.554). The critical
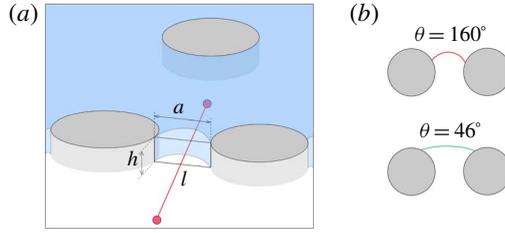
FIGURE 4. (*a*) Pore-scale perspective for the scaling of pressure fluctuations. The diagram shows a typical pore being invaded. The characteristic distance between the pore centres is *l* (red line), the post height is *h*, and a characteristic throat size is *a*. (*b*) Typical configurations of the fluid–fluid interface in drainage and imbibition. Burst events are prevalent in drainage and the typical radius of out-of-plane curvature is of order *a*. Overlap events are prevalent in imbibition and the typical radius of out-of-plane curvature is an order of magnitude greater than *a*.

pressures of touch and overlap are always smaller than the critical pressures of burst events (Cieplak & Robbins 1988, 1990; Primkulov *et al.* 2018), so the magnitude of the pressure fluctuations decreases as the substrate becomes more wetting to the invading fluid (figure 3*b*).

To gain further insight into the difference in the pressure signature between low and high *Ca*, we decompose the injection-pressure signal into its global trend and fluctuating components with block James–Stein wavelet decomposition (Cai 2002) (see figure 3*c*,*d*). We compute the standard deviation of the fluctuating component of the pressure signal for both drainage and imbibition conditions ($\theta = 160°$ and $46°$, respectively) for a wide range of *Ca*, and find that it exhibits two distinct regimes (figure 3*e*). At low *Ca*, pressure fluctuations are controlled by the stick-slip-type changes in local Laplace pressures. In contrast, at high *Ca*, pressure fluctuations are controlled by changes in the effective hydraulic conductance of the dominant flow channels. In the limit of high *Ca*, the Laplace pressure drop is negligible in comparison with the viscous pressure gradient, but the dominant flow channels are rearranged slightly as the fingers grow (see movie S2 in the supplementary materials). Since the pore geometry has a heterogeneous distribution of throat sizes, shifts in the dominant flow channels result in viscosity-driven pressure fluctuations at high *Ca*.

Scaling arguments support the findings from the model simulations. Let us take a pore-scale perspective (see figure 4). Invading a single pore involves overcoming a capillary pressure and pushing defending fluid out through a throat of width *a* and height *h* at a speed proportional to the injection rate. The capillary pressure is $p_{cap} \approx \gamma((1/h) + (1/af(\theta)))$, where $f(\theta)$ is a wettability-dependent function that takes a value ~ 1 near drainage and ~ 10 near strong imbibition (figure 4*b*). Taking variations of p_{cap} with *a* yields

$$\delta p_{cap} \sim \frac{\gamma}{a^2 f(\theta)} \delta a. \qquad (4.1)$$

The characteristic flow velocity through a typical throat is $u = (k(a, h)/\mu)(p_{visc}/l)$, where $k(a, h) = R_h^2/8$ is the rectangular channel permeability and $R_h = ah/2(a + h)$ the hydraulic radius. Thus the viscous pressure is $p_{visc} \sim (32(a + h)^2/a^2 h^2)\mu u l = 32\mu u l/h^2(1 + h/a)^2$. Taking variations of p_{visc} with *a* yields

$$\delta p_{visc} \sim \frac{64\mu u l}{h^2}(1 + h/a)\frac{h}{a^2}\delta a = \frac{64\mu u l}{h a^2}(1 + h/a)\delta a. \qquad (4.2)$$

875 R4-9

The magnitude of the total characteristic pressure fluctuation is $\delta p_{cap} + \delta p_{visc}$, and its two components are comparable when $\delta p_{visc}/\delta p_{cap} \sim 1$. Using (4.1) and (4.2),

$$\frac{\delta p_{visc}}{\delta p_{cap}} \sim \frac{64\mu u l}{ha^2}(1+h/a)\frac{a^2 f(\theta)}{\gamma} = Ca f(\theta) 64 \frac{l}{h}(1+h/a) \sim 1, \qquad (4.3)$$

which implies a crossover Ca,

$$Ca^* \sim \frac{h}{64 f(\theta)(1+h/a)l}, \qquad (4.4)$$

between flow-rate-independent and flow-rate-dependent pressure fluctuations (figure 3*e*). The above argument suggests two interesting implications. First, one can potentially infer the characteristic pore size of the material from the fluctuations of the pressure signal in both viscous-dominated and capillary-dominated flow regimes. This is especially useful when visualization of the flow in pore space is not possible, which is the case in most porous materials. Second, the characteristic h, a and l used in this study yield $Ca^* \approx 10^{-3}/f(\theta)$, which reduces to $Ca^* \sim 10^{-3}$ for drainage and $Ca^* \sim 10^{-4}$ for imbibition, in agreement with the data in figure 3(*e*). This means that one should expect the transition from capillary-dominated to viscous-dominated flow regimes at different Ca^* in drainage and imbibition. The order of magnitude of $f(\theta)$ was obtained by calculating Δp_{crit} for all pore throats at $\theta \in \{46°, 160°\}$ with the quasi-static model (Primkulov *et al.* 2018) and taking an average of $f(\theta) = \gamma/a\Delta p_{crit}$ for each contact angle. Finally, the viscous pressure fluctuation component scales as $\delta p_{visc} \sim \mu u$, which is equivalent to $\delta p_{visc} \sim Ca$ when interfacial tension is kept constant. This explains the slope of the viscous-dominated portion of the graph in figure 3(*e*).

5. Conclusion

Overall, our moving-capacitor network model provides new fundamental insights into the dynamics of immiscible fluid–fluid displacement in porous media for a wide range of Ca and wettabilities. The model completes the picture of the displacement by covering both high and low Ca, which allows one, for the first time, to reproduce experimental observations of invading fluid patterns (Zhao *et al.* 2016), injection pressure and front velocity in drainage (Måløy *et al.* 1992; Furuberg *et al.* 1996; Moebius & Or 2012) and imbibition. Our observations and scaling arguments on the transition from a viscous-dominated to a capillary-dominated flow regime suggest that it is possible to infer the character of the multiphase-flow displacement purely from the injection-pressure signal. This poses an exciting prospect for detailed experiments.

Supplementary movies

Supplementary movies are available at https://doi.org/10.1017/jfm.2019.554.

References

AKER, E., MÅLØY, K. J. & HANSEN, A. 1998*a* Simulating temporal evolution of pressure in two-phase flow in porous media. *Phys. Rev. E* **58** (2), 2217–2226.

AKER, E., MÅLØY, K. J., HANSEN, A. & BATROUNI, G. G. 1998*b* A two-dimensional network simulator for two-phase flow in porous media. *Trans. Porous Med.* **32** (2), 163–186.

AL-GHARBI, M. S. & BLUNT, M. J. 2005 Dynamic network modeling of two-phase drainage in porous media. *Phys. Rev.* E **71** (1), 016308.

ARNÉODO, A., COUDER, Y., GRASSEAU, G., HAKIM, V. & RABAUD, M. 1989 Uncovering the analytical Saffman–Taylor finger in unstable viscous fingering and diffusion-limited aggregation. *Phys. Rev. Lett.* **63** (9), 984–987.

BENSIMON, D., KADANOFF, L. P., LIANG, S., SHRAIMAN, B. I. & TANG, C. 1986 Viscous flows in two dimensions. *Rev. Mod. Phys.* **58** (4), 977–999.

BERG, S., OTT, H., KLAPP, S. A., SCHWING, A., NEITELER, R., BRUSSEE, N., MAKURAT, A., LEU, L., ENZMANN, F., SCHWARZ, J.-O. *et al.* 2013 Real-time 3D imaging of Haines jumps in porous media flow. *Proc. Natl Acad. Sci. USA* **110** (10), 3755–3759.

BIROVLJEV, A., FURUBERG, L., FEDER, J., JØSSANG, T., MÅLØY, K. J. & AHARONY, A. 1991 Gravity invasion percolation in two dimensions – experiment and simulation. *Phys. Rev. Lett.* **67**, 584–587.

BISCHOFBERGER, I., RAMACHANDRAN, R. & NAGEL, S. R. 2015 An island of stability in a sea of fingers: emergent global features of the viscous-flow instability. *Soft Matt.* **11** (37), 7428–7432.

BLUNT, M. J. 2001 Flow in porous media pore-network models and multiphase flow. *Curr. Opin. Colloid Interface Sci.* **6** (3), 197–207.

CAI, T. T. 2002 On block thresholding in wavelet regression: adaptivity, block size, and threshold level. *Statistica Sin.* **12**, 1241–1273.

CHANDLER, R., KOPLIK, J., LERMAN, K. & WILLEMSEN, J. F. 1982 Capillary displacement and percolation in porous media. *J. Fluid Mech.* **119**, 249–267.

CHEN, J. D. 1987 Radial viscous fingering patterns in Hele-Shaw cells. *Exp. Fluids* **5** (6), 363–371.

CHEN, J. D. & WILKINSON, D. 1985 Pore-scale viscous fingering in porous media. *Phys. Rev. Lett.* **55** (18), 1892–1895.

CIEPLAK, M. & ROBBINS, M. O. 1988 Dynamical transition in quasistatic fluid invasion in porous media. *Phys. Rev. Lett.* **60** (20), 2042–2045.

CIEPLAK, M. & ROBBINS, M. O. 1990 Influence of contact angle on quasistatic fluid invasion of porous media. *Phys. Rev.* B **41** (16), 11508–11521.

CONTI, M. & MARCONI, U. M. B. 2010 Diffusion limited propagation of burning fronts. In *WIT Transactions on Ecology and the Environment*, vol. 137, pp. 37–45. WIT Press.

DACCORD, G. G., NITTMANN, J. & STANLEY, H. E. 1986 Radial viscous fingers and diffusion-limited aggregation: fractal dimension and growth sites. *Phys. Rev. Lett.* **56** (4), 336–339.

FERER, M., JI, C., BROMHAL, G. S., COOK, J., AHMADI, G. & SMITH, D. H. 2004 Crossover from capillary fingering to viscous fingering for immiscible unstable flow: experiment and modeling. *Phys. Rev.* E **70** (1), 016303.

FERNÁNDEZ, J. F., ALBARRÁN, J. M., FERNANDEZ, J. F. & ALBARRAN, J. M. 1990 Diffusion-limited aggregation with surface tension: scaling of viscous fingering. *Phys. Rev. Lett.* **64** (18), 2133–2136.

FERNANDEZ, J. F., RANGEL, R. & RIVERO, J. 1991 Crossover length from invasion percolation to diffusion-limited aggregation in porous media. *Phys. Rev. Lett.* **67** (21), 2958–2961.

FRETTE, V., FEDER, J., JØSSANG, T. & MEAKIN, P. 1992 Buoyancy-driven fluid migration in porous media. *Phys. Rev. Lett.* **68**, 3164–3167.

FURUBERG, L., MÅLØY, K. J. & FEDER, J. 1996 Intermittent behavior in slow drainage. *Phys. Rev.* E **53** (1), 966–977.

GJENNESTAD, M. A., VASSVIK, M., KJELSTRUP, S. & HANSEN, A. 2018 Stable and efficient time integration of a dynamic pore network model for two-phase flow in porous media. *Frontiers Phys.* **6**, 56.

HAINES, W. B. 1930 Studies in the physical properties of soil. V. The hysteresis effect in capillary properties, and the modes of moisture distribution associated therewith. *J. Agri. Sci.* **20** (1), 97–116.

HOFFMAN, R. L. 1975 A study of the advancing interface. I. Interface shape in liquid–gas systems. *J. Colloid Interface Sci.* **50** (2), 228–241.

HOLTZMAN, R. & JUANES, R. 2010 Crossover from fingering to fracturing in deformable disordered media. *Phys. Rev.* E **82** (4), 046305.

HOLTZMAN, R. & SEGRE, E. 2015 Wettability stabilizes fluid invasion into porous media via nonlocal, cooperative pore filling. *Phys. Rev. Lett.* **115** (16), 164501.

HOLTZMAN, R., SZULCZEWSKI, M. L. & JUANES, R. 2012 Capillary fracturing in granular media. *Phys. Rev. Lett.* **108** (26), 264504.

HOMSY, G. M. 1987 Viscous fingering in porous media. *Annu. Rev. Fluid Mech.* **19** (1), 271–311.

JOEKAR-NIASAR, V. & HASSANIZADEH, S. M. 2012 Analysis of fundamentals of two-phase flow in porous media using dynamic pore-network models: a review. *Crit. Rev. Environ. Sci. Technol.* **42** (18), 1895–1976.

JOEKAR-NIASAR, V., HASSANIZADEH, S. M. & DAHLE, H. K. 2010 Non-equilibrium effects in capillarity and interfacial area in two-phase flow: dynamic pore-network modelling. *J. Fluid Mech.* **655**, 38–71.

JUNG, M., BRINKMANN, M., SEEMANN, R., HILLER, T., SANCHEZ DE LA LAMA, M. & HERMINGHAUS, S. 2016 Wettability controls slow immiscible displacement through local interfacial instabilities. *Phys. Rev. Fluids* **1** (7), 074202.

KADANOFF, L. P. 1985 Simulating hydrodynamics: a pedestrian model. *J. Stat. Phys.* **39** (3–4), 267–283.

KNUDSEN, H. A. & HANSEN, A. 2002 Relation between pressure and fractional flow in two-phase flow in porous media. *Phys. Rev.* E **65** (5), 056310.

LEE, H., GUPTA, A., HATTON, T. A. & DOYLE, P. S. 2017 Creating isolated liquid compartments using photopatterned obstacles in microfluidics. *Phys. Rev.* A **7** (4), 044013.

LENORMAND, R., TOUBOUL, E. & ZARCONE, C. 1988 Numerical models and experiments on immiscible displacements in porous media. *J. Fluid Mech.* **189**, 165–187.

LENORMAND, R. & ZARCONE, C. 1985 Invasion percolation in an etched network: measurement of a fractal dimension. *Phys. Rev. Lett.* **54** (20), 2226–2229.

LENORMAND, R., ZARCONE, C. & SARR, A. 1983 Mechanisms of the displacement of one fluid by another in a network of capillary ducts. *J. Fluid Mech.* **135**, 337–353.

LI, S., LOWENGRUB, J. S., FONTANA, J. & PALFFY-MUHORAY, P. 2009 Control of viscous fingering patterns in a radial Hele-Shaw cell. *Phys. Rev. Lett.* **102** (17), 174501.

MÅLØY, K. J., FEDER, J. & JØSSANG, T. 1985 Viscous fingering fractals in porous media. *Phys. Rev. Lett.* **55** (24), 2688–2691.

MÅLØY, K. J., FURUBERG, L., FEDER, J. & JOSSANG, T. 1992 Dynamics of slow drainage in porous media. *Phys. Rev. Lett.* **68** (14), 2161–2164.

MEAKIN, P., FEDER, J., FRETTE, V. & JØSSANG, T. 1992 Invasion percolation in a destabilizing gradient. *Phys. Rev.* A **46** (6), 3357–3368.

MEAKIN, P. & TARTAKOVSKY, A. M. 2009 Modeling and simulation of pore-scale multiphase fluid flow and reactive transport in fractured and porous media. *Rev. Geophys.* **47** (3), RG3002.

MEAKIN, P., TOLMAN, S. & BLUMEN, A. 1989 Diffusion-limited aggregation. *Proc. R. Soc. Lond.* A **423** (1864), 133–148.

MOEBIUS, F. & OR, D. 2012 Interfacial jumps and pressure bursts during fluid displacement in interacting irregular capillaries. *J. Colloid Interface Sci.* **377** (1), 406–415.

NIEMEYER, L., PIETRONERO, L. & WIESMANN, H. J. 1984 Fractal dimension of dielectric breakdown. *Phys. Rev. Lett.* **52** (12), 1033–1036.

NITTMANN, J., DACCORD, G. & STANLEY, H. E. 1985 Fractal growth viscous fingers: quantitative characterization of a fluid instability phenomenon. *Nature* **314** (6007), 141–144.

ODIER, C., LEVACHÉ, B., SANTANACH-CARRERAS, E. & BARTOLO, D. 2017 Forced imbibition in porous media: a fourfold scenario. *Phys. Rev. Lett.* **119** (20), 208005.

PATERSON, L. 1981 Radial fingering in a Hele Shaw cell. *J. Fluid Mech.* **113**, 513–529.

PRIMKULOV, B. K., TALMAN, S., KHALEGHI, K., RANGRIZ SHOKRI, A., CHALATURNYK, R., ZHAO, B., MACMINN, C. W. & JUANES, R. 2018 Quasistatic fluid–fluid displacement in porous media: invasion–percolation through a wetting transition. *Phys. Rev. Fluids* **3**, 104001.

RABBANI, H. S., ZHAO, B., JUANES, R. & SHOKRI, N. 2018 Pore geometry control of apparent wetting in porous media. *Sci. Rep.* **8** (1), 15729.

SAFFMAN, P. G. & TAYLOR, G. 1958 The penetration of a fluid into a porous medium or Hele-Shaw cell containing a more viscous liquid. *Proc. R. Soc. Lond.* A **245** (1242), 312–329.

STOKES, J. P., WEITZ, D. A., GOLLUB, J. P., DOUGHERTY, A., ROBBINS, M. O., CHAIKIN, P. M. & LINDSAY, H. M. 1986 Interfacial stability of immiscible displacement in a porous medium. *Phys. Rev. Lett.* **57** (14), 1718–1721.

STRANG, G. 2007 *Computational Science and Engineering*. Wellesley–Cambridge Press.

SYGOUNI, V., TSAKIROGLOU, C. D. & PAYATAKES, A. C. 2006 Capillary pressure spectrometry: toward a new method for the measurement of the fractional wettability of porous media. *Phys. Fluids* **18** (5), 053302.

SYGOUNI, V., TSAKIROGLOU, C. D. & PAYATAKES, A. C. 2007 Using wavelets to characterize the wettability of porous materials. *Phys. Rev.* E **76** (5), 056304.

TOUSSAINT, R., LØVOLL, G., MÉHEUST, Y., MÅLØY, K. J. & SCHMITTBUHL, J. 2005 Influence of pore-scale disorder on viscous fingering during drainage. *Europhys. Lett.* **71** (4), 583–589.

TROJER, M., SZULCZEWSKI, M. L. & JUANES, R. 2015 Stabilizing fluid–fluid displacements in porous media through wettability alteration. *Phys. Rev. Appl.* **3** (5), 054008.

TRYGGVASON, G. & AREF, H. 1983 Numerical experiments on Hele Shaw flow with a sharp interface. *J. Fluid Mech.* **136**, 1–30.

WILKINSON, D. 1984 Percolation model of immiscible displacement in the presence of buoyancy forces. *Phys. Rev.* A **30** (1), 520–531.

WILKINSON, D. & WILLEMSEN, J. F. 1983 Invasion percolation: a new form of percolation theory. *J. Phys.* A **16** (14), 3365–3376.

WITTEN, T. A., SANDER, L. M. & SANDER, I. M. 1981 Diffusion-limited aggregation, a kinetic critical phenomenon. *Phys. Rev. Lett.* **47** (19), 1400–1403.

YORTSOS, Y. C., XU, B. & SALIN, D. 1997 Phase diagram of fully-developed drainage in porous media. *Phys. Rev. Lett.* **79** (23), 4581–4584.

ZHAO, B., MACMINN, C. W. & JUANES, R. 2016 Wettability control on multiphase flow in patterned microfluidics. *Proc. Natl Acad. Sci. USA* **113** (37), 10251–10256.